2009 SOLAS Summer School

Introduction to Marine Aerosols



Eric S. Saltzman Earth System Science Univ. of CA, Irvine

What is driving marine aerosol research?

- geochemical cycles
 - global elemental fluxes, transport properties
- acidification (sulfur, nitrogen)
- climate change
 - direct/indirect effects
 - aerosol optical properties, aerosol/cloud interactions
 - interactions between aerosols/oxidation capacity
- nutrients
 - N deposition
 - budgets in coastal, oligotrophic regions
 - desert dust and iron deposition, HNLC regions
- human health
 - air quality, airborne pathogen transport

Earth's energy balance...



Aerosol effects on Earth's radiation budget

global radiative forcing (2000 vs 1750)



Direct effects (cloud-free): ●scatter → cooling ●absorption → heating

Indirect effects (clouds):
•more (but smaller) droplets → scatter (Twomey)
•more droplets → longer cloud lifetime (Albrecht)
•absorption → heating → evaporates clouds

Some questions...

•What are marine aerosols made of ?

•Where do they come from?

•How long do they stay in the atmosphere? How are they removed?

•How do they evolve while in the atmosphere?

How do they interact with the climate system?
forcing
feedback

 How will they change in the future as a result of industrialization and climate change?

•...what don't we know about marine aerosols that we need to know?

Things we "know we know"... (or don't)...

"...there are known knowns. These are things we know that we know. We also know there are known unknowns; that is to say we know there are some things we do not know. But there are also unknown unknowns. These are things we don't know we don't know."

Donald Rumsfeld

"Unknown knowns" ...things we know that might not be true Eric Saltzman

- "nucleation is rare"
- "sulfate is the source of CCN over the oceans"
- "sea spray doesn't make submicron aerosols"
- "organics in marine aerosols are derived from land"
- "there's not enough iodine out there to matter..."

Terminology...

• aerosol - a dispersion of solid and liquid particles suspended in gas (air).

note: "aerosol" is defined as the dispersion of both particles and gas, but in common practice it is used to refer to the particles only!

• primary aerosol - emitted directly into the atmosphere.

soot, fly ashSaharan dust, pollensea spray

•<u>secondary aerosol</u> - created by nucleation of new particles, aggregation of existing particles, or growth of preexisting particles from gas phase molecules (gas to particle conversion).

either type may be natural or anthropogenic or both

•*internal vs. external mixtures* – in an internally mixed aerosol all particles have the same composition

How much aerosol is there? typically ~10's of ug/m³ (air ~1kg/m³)

The aerosol size distribution...



The log-normal aerosol size distribution...



Diameter (µm)

aerosol removal and lifetimes...







diffusion, coagulation

Humidity and aerosol size ...

 hygroscopic aerosols grow/shrink with RH (with hysteresis)





 aerosol size strongly affects light scattering cross-section

The aerosol modes...

- Aitken mode 0.01-0.1 um
- accumulation mode 0.1-1 um
- coarse mode >1 um

and sometimes, the elusive

• nucleation mode <0.01 um



Cloud processing



- cloud residue mode due to evaporation of cloud droplets
- "splits" Aitken mode in marine air.

a process-oriented view of the size distribution ... competition between production/transformation/removal

aerosol modes:



Observed marine aerosol number distributions (>10,000!)



The chemical perspective ... a chemical size distribution

- chemical size distributions resemble mass, not number
- sulfate and organics dominate the accumulation mode, but there's a surprising amount of seasalt
- there are <u>a lot</u> of unidentified organics
- the coarse mode has the expected mechanically generated aerosols, but also nitrate and sometimes sulfate



Marine aerosol system...



soil dust (aluminosilicates, trace metals, organics) biomass burning (organics) pollutant aerosols (sulfate, nitrate, ammonium, soot) and precursor gases (50₂, NH₃, NO₂, VOC, OVOC)



air/sea gas exchange sea spray bubble bursting

DMS, VOC's, OVOC's halocarbons

iodine cmpds? non-volatile organics? skeletal fragments...?

Dust (mineral aerosols)
diameter size: 2-300 μm
main material: sand, silt, clay
includes essential trace metals such as Fe
consists of insoluble and soluble fractions

Mineral Dust



Seasalt aerosols...

wind \rightarrow bubbles \rightarrow spray

whitecap coverage $W \alpha U^{3+}!$



seasalt production via bubble bursting...

- film drops (many, small, organics)
- jet drops (fewer, larger)



Seasalt aerosols...



Seasalt – number and mass as a function of surface wind speed



 10^{3} (_{r-}u 6rt)/W 195 ger & Garre et al. orvath et al., ulkarni et al., 16 n & Lepple, 1985 & Tsunoggi, 1988, et gl., 1989 1990, 12&18.3 m, - å McGovern et al., 1994 McKay et al., 1994 Chabas & Lefevre, 2000 NAM, 1983, k NOVAM, 1993, j O'Dowd et al., 1997a, m Quadrafic, n Linear, o 100 20 0 10 30 $U_{10}/(m \ s^{-1})$

seasalt particle number

seasalt mass

10⁴

(compiled by Lewis and Schwartz, 2004)

Marine aerosol mass fractions...



Fraction of submicron aerosol derived from nss sulfate, sea salt and "residual"

Quinn et al., 2000

The sulfur story (in brief) ...

- emissions: fossil fuel SO₂, volcanic SO₂, oceanic DMS
- DMS oxidation ... gas phase ... complex!



Happily, gas phase SO_2 oxidation is simple...

$$SO_{2} + OH \xrightarrow{M} HOSO_{2}$$
$$HOSO_{2} + O_{2} \longrightarrow HO_{2} + SO_{3}$$
$$SO_{3} + H_{2}O + H_{2}O \longrightarrow H_{2}SO_{4} + H_{2}O$$

sadly, most SO₂ oxidation occurs in the aqueous phase... some basics... $SO_2(\alpha) \longleftrightarrow SO_2(\alpha)$

$$SO_{2} + H_{2}O \xrightarrow{pK_{1} \sim 4} HSO_{3}^{-} + H^{+}$$
$$HSO_{3}^{-} + H_{2}O \xrightarrow{pK_{2} \sim 8} SO_{3}^{-} + H^{+}$$







FIGURE 8.8 Range of expected aqueous S(IV) concentrations as a function of acidity for gas-phase SO_2 concentrations of 0.2–200 ppb (adapted from Martin, 1984).

heterogeneous oxidation of SO_2

- in-cloud oxidation
 - weakly buffered, pH ~4
 - oxidation by H_2O_2
 - aerosol growth, split Aitken mode



- seasalt aerosols
 - strongly buffered by carbonate system
 - rapid oxidation by O_3
 - then slower oxidation by H_2O_2 (also OH, halogen radicals...)
 - rapid loss by deposition

(Chameides and Stelson, 1992)

Organic aerosols - burning

soot – "elemental carbon" formed in flames little spectral dependence carbon-only





*"brown carbon":*sugars
alcohols
aromatics (and methoxy)
di/tri acids
ketoacids
hydroxyacids

Marine-derived organic aerosols...





(O'Dowd et al., 2004)

More organics in marine aerosols ...

SEM micrograph of an Arctic marine aerosol sample. Single particles (diameter 1 μ m) and groups of particles are lying between the fibers of the filter. About half of the particles are coated by an organic layer.



Negative TOF-SIMS spectra of marine aerosol.

mass $255 \rightarrow C_{15}H_{31}COOH$ (palmitic acid)



Tervahattu et al., 2002

Iodine and nucleation...

- coastal I_2 emissions $\rightarrow I_2O_5$
- macroalgae
- open ocean CH_2I_2 , I_2 ...?
- broader importance?







O'Dowd et al., 2002a,b

The single-particle view...





Murphy et al., 1998; Buseck and Posfai, 1999

Global view of aerosol optical depth... MODIS Sept. 2000

fine particles •urban pollution •biomass burning

coarse particles •mineral dust



Aerosol properties from space...

Aersol Optical Thickness

Aerosol size info •angstrom coefficient

(a) May 1997 Aerosol optical depth at 865 nm from Polder on ADEOS



0.0 0.1 0.2 0.3 0.4 0.5 (b) May 1997 Ängström coefficient



-0.2 0.0 0.2 0.4 0.8 0.8 1.0 12

POLOER data: CNES/NASDA; Protessing: LOA/LSCE

Figure 5.3: (a) Aerotol optical depth and (b) Ångstrom exponent from POLDER satellike data for May 1997 (Denné et al., 1999). The largest optical depths over the Alaotic Ocean are from sorth African dust. The Ångstrom exposent expresses the wavelength dependence of scattered light between 670 and 865 nm. The African dust plume has a small Ångstrom exposent due to the importance of oceanse mode aerotols whereas the larger Ångstrom exponenti around the continents show the importance of accumulation mode aerotols in those locations.

POLDER-1 (IPCC, Deuze et al., 1999)

Understanding forcing/feedback requires realistic

aerosols in climate models...

annual average sources (kg km⁻² hr¹)

IPCC, 2001







Modeled atmospheric aerosols...





MODIS (Kaufman et al., 2002)

GOCART model (Chin et al., 2002)

optical thickne

Acrosol

0.00

What's left to do...?

- *in situ* aerosol/cloud radiation experiments
- understand the satellite signals
- model "real" aerosols (shapes, internal/external mixtures)
- marine microlayer
- "life cycle" of marine aerosol organics
- gas phase \rightarrow aerosol chemistry (sulfur, iodine, organics)
- aerosol \rightarrow gas phase cycling (halogens)
- coupled aerosol/chemistry/climate experiments and models
- *marine aerosols are woefully undersampled!*

... the end!!





Plate 2. Satellite observations and model simulations of global aerosol optical thickness (at 0.55 µm) for September, 2000. Left, NASA Moderate Resolution Imaging Spectroradiometer (on EOS) satellite observations of (upper left) fine aerosols and (lower left) total aerosols. Right, results from the GOCART model simulating sources, transport, and deposition of (upper right) anthropogenic sulfate, (middle right) black carbon and particulate organic matter, and (lower right) dust and sea salt. Data obtained from the NASA Giovanni online data system (http://disc.sci.gsfc.nasa.gov/giovanni). See Kaufman et al. [2002] and Chin et al. [2002] for further information.